

RESEARCH ARTICLE

The Early Cretaceous coal-forming plants of southern part of East Siberia and Russian Far East

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Abstract

For the first time the plants that gave rise to the Early Cretaceous coals of Transbaikalia (Khilok, Chita-Ingoda, Bukachacha, and Turga-Kharanor basins), Amur River region (Bureya Basin), and Primorye region (Razdolnaya River and Partizansk River basins) have been obtained. The plants that produced coals mainly belong to the ginkgoaleans (*Pseudotorellia*, *Sphenobaiera*, and *Baierella*), bennettites, and conifers having pinaceous, taxodiaceous, araucariaceous affinity, as well as extinct cheirolepidiaceus and miroviaceous plants. A parallel palynological study has identified a number of the same elements in addition to the cyatheaceous, gleicheniaceus, osmundaceous, and schizaeaceous ferns.

KEYWORDS

Amur River region, coal-forming plants, Cretaceous, Primorye region, Transbaikalia

1 | INTRODUCTION

The coal formation in the geological past was linked with the development of plant communities, which produced a large biomass at the same time possessed relatively simple structure. A considerable part of phytomass was removed from swamp ecosystems and became mortmass. The coal formation involving rapid burial of plant remains has paleoecological and taphonomical significance (e.g. Barbacka et al., 2015; Ferguson et al., 2010; Greb, Dimichele, & Gastaldo, 2006; Iannuzzi, 2010; Jasper, Hartkopf-Fröder, Flajs, & Littke, 2010; Michaelsen, 2002; Philippe et al., 2002; Steart, Collinson, Scott, Glasspool, & Hooker, 2007; Upchurch, 1995; Villalba-Breva, Martín-Closas, Marmi, Gomez, & Fernández-Marrón, 2012).

The coal seams are particularly the subject of paleobotanical research. The study of plant material has great importance for knowledge of the genesis of coal, its composition, and its quality.

Many elements of the coal-forming plant community can still be identified in the coal. At the time of the mire formation and growth, climatic conditions were not suitable for sediment transport into the basin particularly biogenic components, with almost complete absence of clastic sediments. Therefore, it is possible to assume that the plants that gave rise to the coal deposits were autochthonous and provided the first elements of peat accumulation. Since clastic deposits represent the material transported from the source area, phytofossil assemblages from the terrigenous deposits between the

coal layers are often mixed, consisting of elements from both the slope and lowland communities.

In the present study, coal-bearing deposits of Transbaikalia (Chita-Ingoda, Khilok, Bukachacha, Turga-Kharanor basins), Amur River (Bureya Basin) and Primorye regions (Razdolnaya and Partizansk basins) were studied (Figure 1).

These deposits are of Tithonian–Valanginian, and Barremian–Albian stages of coal formation in these areas, of which, the Barremian–Albian stage occupied a large part (Figure 2). Coal accumulation was manifested on vast areas of Siberia and the Far East. During the Barremian–Aptian, the main coal-forming plants varied in composition depending on environments.

During this period, favorable conditions for the existence of mire vegetation occurred. The abundant plant material was feedstock for peat accumulation and coal origin. Coal-forming plants were buried in close proximity to their habitats. The extremely thin cuticles of ferns desintegrate after chemical maceration of coal, but their spores, having exine resistant to acids and alkali, provide evidence of the existence of ferns in mire vegetation (Taylor, Taylor, & Krings, 2009). Both palaeobotanical and palynological data reveal that these plants contributed to the coal formation. Usually coal and coaly rocks contain highly fragmented plant material as spores, pollen, and dispersed cuticles. The identifiable latter can be a very significant proxy for determination of taxonomical composition of coal-forming plants, paleoecological

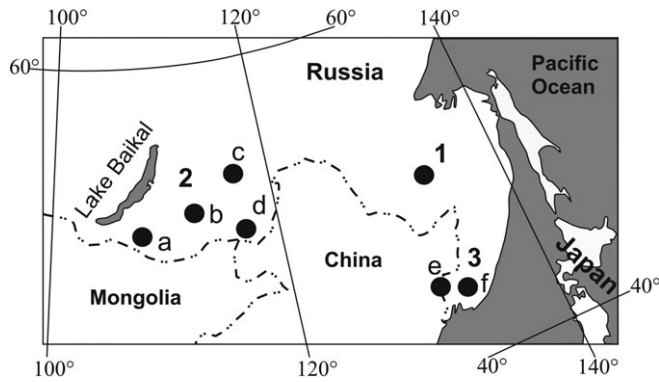


FIGURE 1 The studied Upper Jurassic–Lower Cretaceous coal basins. 1, Amur River region, Bureya Basin; 2, Transbaikalia; 3, Primorye region; a, Khilok Basin; b, Chita-Ingoda Basin; c, Bukachacha Basin; d, Turga-Kharanor Basin; e, Razdolnaya River Basin; f, Partizansk Basin

interpretation of living conditions of swamp plant communities, paleoenvironments, and stratigraphic correlation.

The study of dispersed cuticles began in the 20th Century and was linked to the study of coal geology and emphasized facies

analysis and paleoecology (Kovach & Dilcher, 1984; Litke, 1966; Miner, 1935; Peters, 1963; Roselt & Schneider, 1969; Schneider, 1969; Upchurch, 1995; Upchurch & Askin, 1989). Special studies devote to plant cuticle structures and function (e.g. Barclay, McElwain, Dilcher, & Sageman, 2007; Cutler, 1982; Martin & Juniper, 1970; McElwain & Chaloner, 1996; Stace, 1965). The dispersed cuticles often having preserved systematic characters can be compared readily with the cuticular anatomy of megafossils.

Thereby, our main objective is a recovery of palynomorphs and dispersed cuticles from coals of the basins of the Transbaikalia, Amur River and Primorye regions, definition of their taxonomical composition, and after that to reveal a composition of coal-forming plants.

From the late Albian to the end of the Cretaceous peat accumulation was drastically reduced, especially in the south of Primorye region, where red-bed facies appeared for the first time (Khanchuk, Ratkin, Ryazantseva, Golozubov, & Gonokhova, 1995; Krassilov, 1985). These geological events may have been caused by climatic changes associated with a rainshadow effect of the rising volcanic ranges (Krassilov, 1967, 1985, 1992).

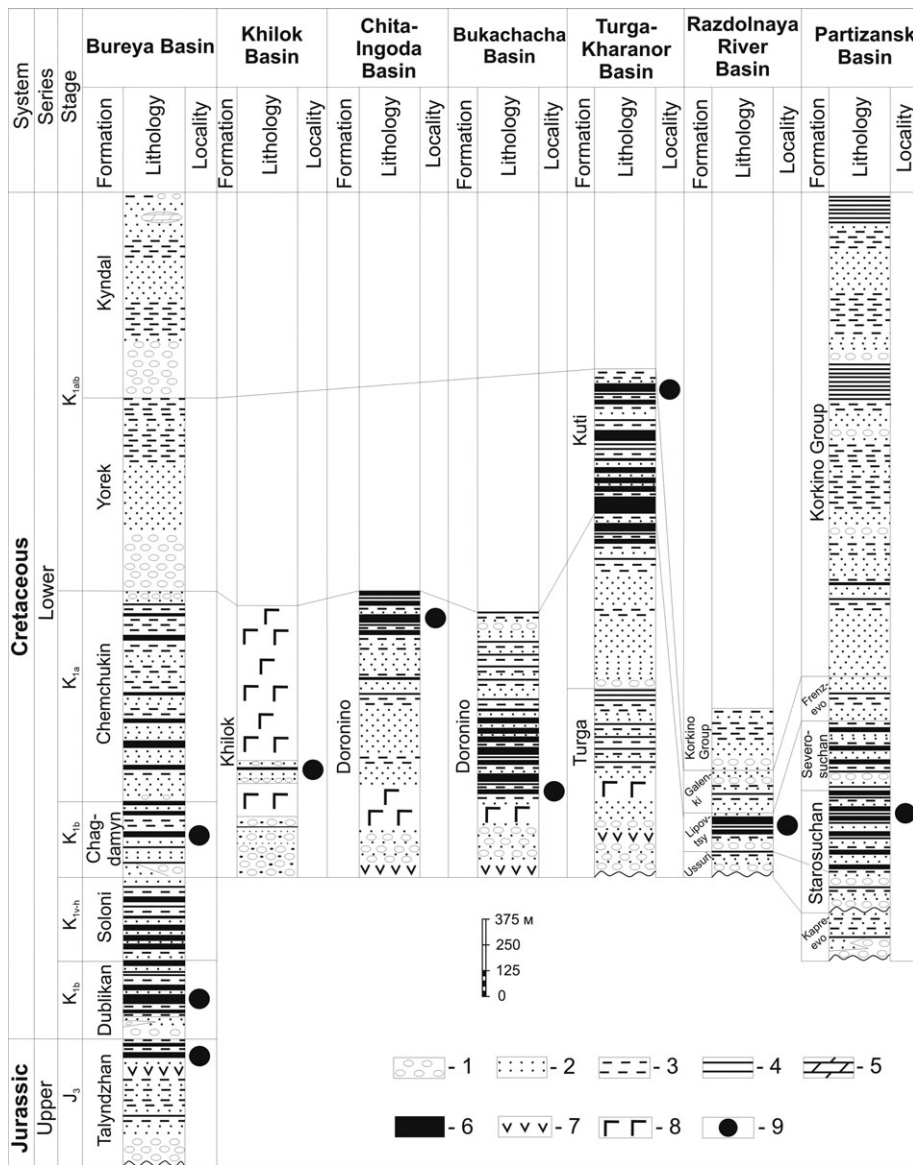


FIGURE 2 The sequences of coal-bearing basins of Amur River region, Transbaikalia, and Primorye region. 1, conglomerate; 2, sandstone; 3, siltstone; 4, mudstone; 5, marl; 6, coal; 7, basalt; 8, andesite; 9, place of coal sampling

The coals are often underlain by rooted tuffaceous sediments (Figure 3) and are overlain by tuffs (e.g. coals of the Talyndzhan Formation in the Bureya Basin and the Lipovtsy Formation in the Razdolnaya Basin; Figure 4). Peat growth was usually terminated by heavy ash falls that buried the swamps, preventing accumulation of thick coal seams (Kovaleva, Markevich, Bugdaeva, Volynets, & Afonin, 2016; Krassilov, 1992). The swamp plant communities and coal accumulation seem sensitive to major environmental crises.

2 | MATERIALS AND METHODS

Abundant plant material was collected from the Lower Cretaceous coal-bearing deposits of Transbaikalia (the Khilok Formation in the

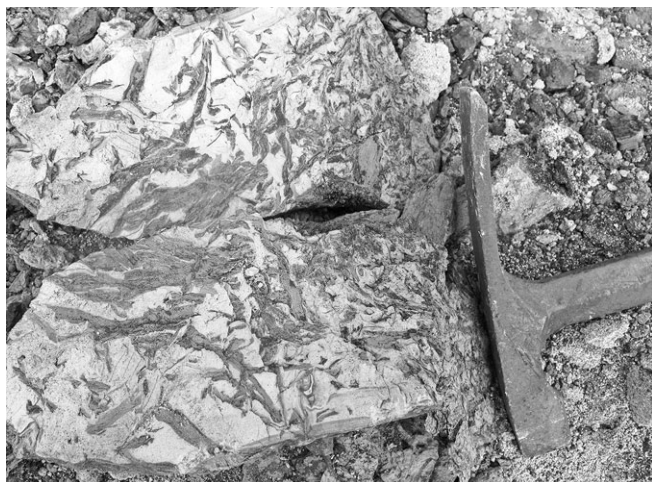


FIGURE 3 The rooted tuffaceous sediments below the coal seam. Lipovtsy Formation, Porechye coal mine, Razdolnaya River Basin. The length of this steel head of the geological hammer is 14.5 cm

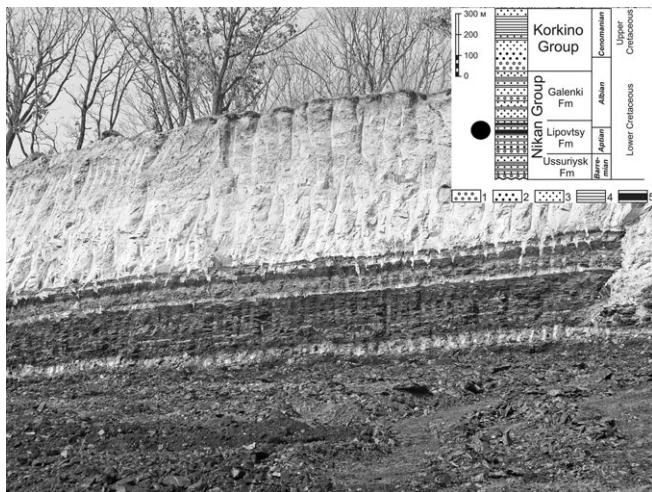


FIGURE 4 Tuff overlying the coal seam. Lipovtsy Formation, Porechye coal mine, Razdolnaya River Basin. Inset shows a schematic sequence of the Cretaceous deposits of the Razdolnaya River Basin. 1, conglomerate, gravelstone; 2, coarse-grained sandstone; 3, medium- and fine-grained sandstone; 4, siltstone; 5, coaly mudstone and coal. The black circle indicates the stratigraphic position of the outcrop

Khilok Basin, the Doronino Formation in the Chita-Ingoda and Bukachacha basins, the Kuti Formation in the Turga-Kharanor Basin), the Upper Jurassic–Lower Cretaceous coal-bearing deposits of the Amur River region (the Talyndzhan, Dublikan, Chagdamyn, and Chemchukin formations in the Bureya Basin), and from the Lower Cretaceous of Primorye region (the Lipovtsy Formation in the Razdolnaya Basin and the Starosuchan Formation in the Partizansk Basin).

Both plant megafossils and coals were collected to obtain dispersed cuticles of the coal-forming plants by bulk maceration. The coals were oxidized by concentrated nitric acid, washed with distilled water, then, treated with 10 % alkali (KOH) and rinsed. The cuticles were mounted on permanent slides for observation under a light microscope (LM) or were placed on standard stubs, splutter coated with gold for scanning electron microscope (SEM) studies. Photography was carried out using a Zeiss Axioscop 40 light microscope with camera AxioCam HRC and a Zeiss EVO-40 scanning electron microscope (IBSS FEB RAS).

3 | RESULTS

3.1 | Amur River region

3.1.1 | Bureya Basin

The Bureya Basin is located in the upper part of the Bureya River (Figure 5) and has been the best subject for the study of the Upper Jurassic to Lower Cretaceous coal-bearing deposits, because the sequence of

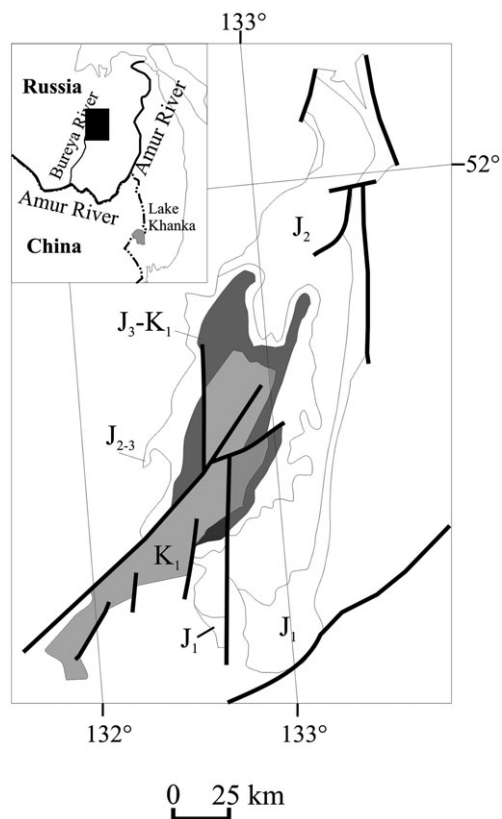


FIGURE 5 Schematic geological map of the Bureya Basin: J_1 , J_2 , J_{2-3} , marine deposits (white color); J_3 - K_1 , transitional brackish-terrestrial deposits with thin coal seams (grey color); K_1 , terrestrial deposits with thick productive coal seams (light-grey color). Thick lines indicate faults

this tectonic structure includes productive coals of the Tithonian, Berriasian, Barremian, and Aptian ages (see Figure 2). The coals of other ages are thin and do not have commercial significance.

The Upper Jurassic strata were deposited in brackish coastal environments. The characteristic features of the Late Jurassic epoch are the homogeneous composition, not clearly defined differentiation, low diversity and smoothed zonation of vegetation. The main coal-forming plants were cyatheaceous ferns, ginkgoaleans, and conifers (Markevich & Bugdaeva, 2014). This is most conspicuous in the Bureya Basin. The coal-bearing Upper Jurassic (Talyndzhan Formation) to Lower Cretaceous (the Dublikan, Soloni, Chagdamyn, and Chemchuko formations) deposits have thickness of about 200–2000 m (see Figure 2).

The palynological assemblage of the Talyndzhan Formation is characterized by predominance of fern spores, mainly assigned to osmundaceous and cyatheaceous ferns, up to 90 % (Figure 6). The participation of diverse mosses is considerable. Gymnosperms are dominated by pollen closely affiliated to Pinaceae and *Ginkgocycadophytus* (Figure 7). The plant megafossils of this stratigraphic unit are

represented by horsetails, bryophytes, ferns, cycadophytes, ginkgoaleans, czekanowskialeans, and conifers (Krassilov, 1972, 1973, 1978; Vakhrameev & Doludenko, 1961). The burials are dominated by ginkgoaleans and czekanowskialeans, the ferns and cycadophytes are highly represented. The representative of ginkgoaleans (*Pseudotorellia angustifolia* Doludenko) sometimes forms monospecific burials; their cuticle remains are common in coals (Figure 8). Perhaps, this arboreal plant prevailed in mire vegetation, osmundaceous and cyatheaceous ferns were in understory. Cycadophytes are often abundant in the clastic beds, but coals entirely lack remains of these plants. Consequently, cycadophytes were not constituent of mire vegetation.

The Lower Cretaceous (the Berriasian–Valanginian) strata are accumulated in vast swampy lowlands. The peculiarity of palynological assemblage from the Dublikan Formation (the Berriasian) is presence of considerable amount of cyatheaceous and lesser osmundaceous fern spores (up to 84 %). Among gymnosperms, *Classopollis* prevails parallel with conifers (see Figure 7). The floristic changes at the Jurassic–Cretaceous boundary are consistent in the

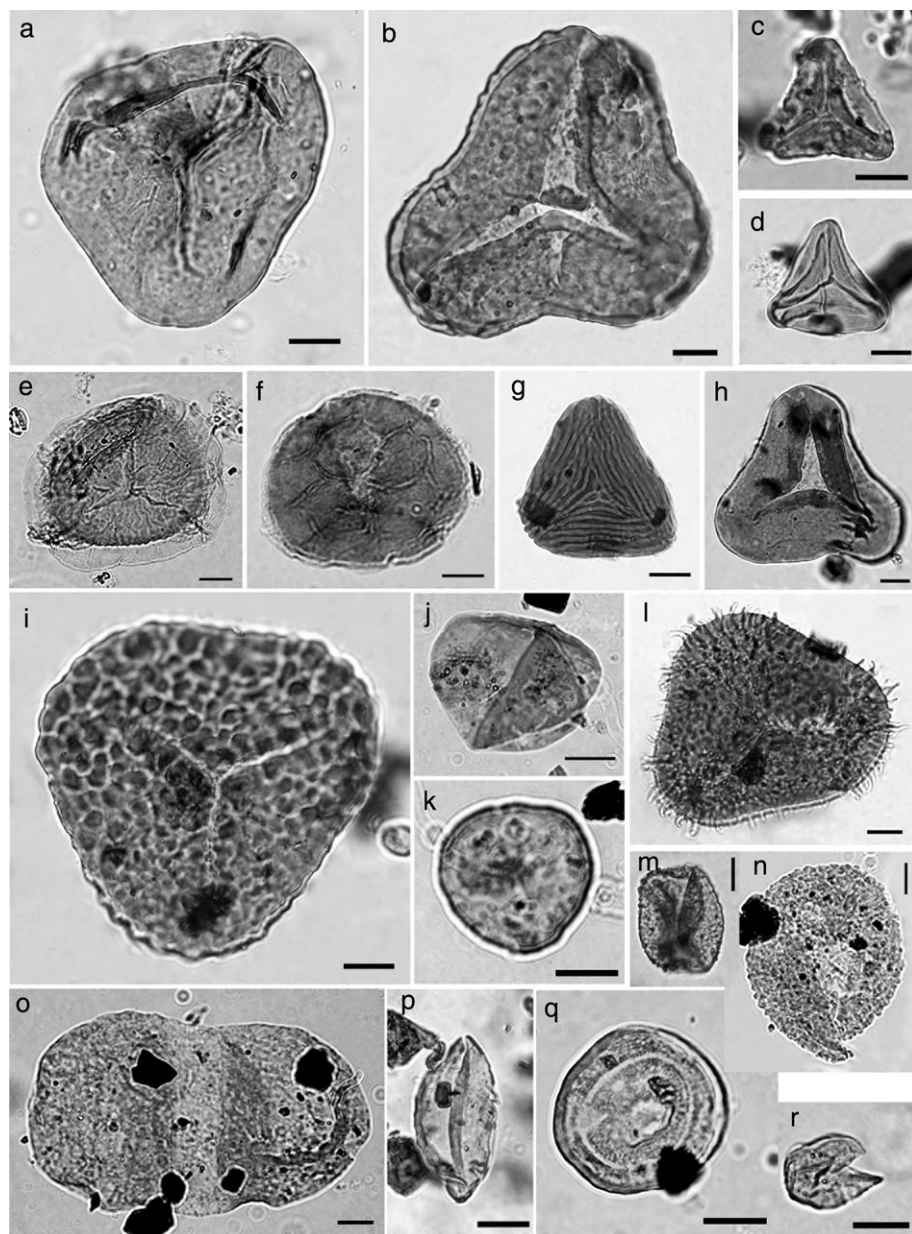


FIGURE 6 The typical spores and pollen of the Lower Cretaceous deposits of Transbaikalia, Amur River region, and Primorye region. Scale bars 10 μ m.

a, *Cyathidites australis* Coup.;
 b, *Impardecispora apiverrucata* (Coup.) Venkat. et Rasa; c, *Gleicheniidites laetus* (Bolch.) Bolch.; d, *Gleicheniidites senonicus* Ross; e, *Denzoisporites vellatus* Weyl. et Krieg.; f, *Rouseisporites reticulatus* Poc.; g, *Cicatricosisporites multicostatus* (Bolch.) Poc.; h, *Concavissimisporites asper* (Bolch.) Poc.; i, *Concavissimisporites variverrucatus* (Coup.) Singh; j, *Cyathidites minor* Coup.; k, *Stereisporites stereoides* (Wills. et Webst.) Dett.; l, *Pilosisporites echinaceus* Bolch.; m, *Osmundacidites nicanicus* (Verb.) E. Sem.; n, *Osmundacidites wellmanii* Coup.; o, *Alisporites bilateralis* Rouse; p, *Ginkgocycadophytus* sp.; q, *Classopollis classoides* Pfl. emend. Poc. et Jans.; r, *Taxodiaceapollenites hiatus* (Pot.) Kremp

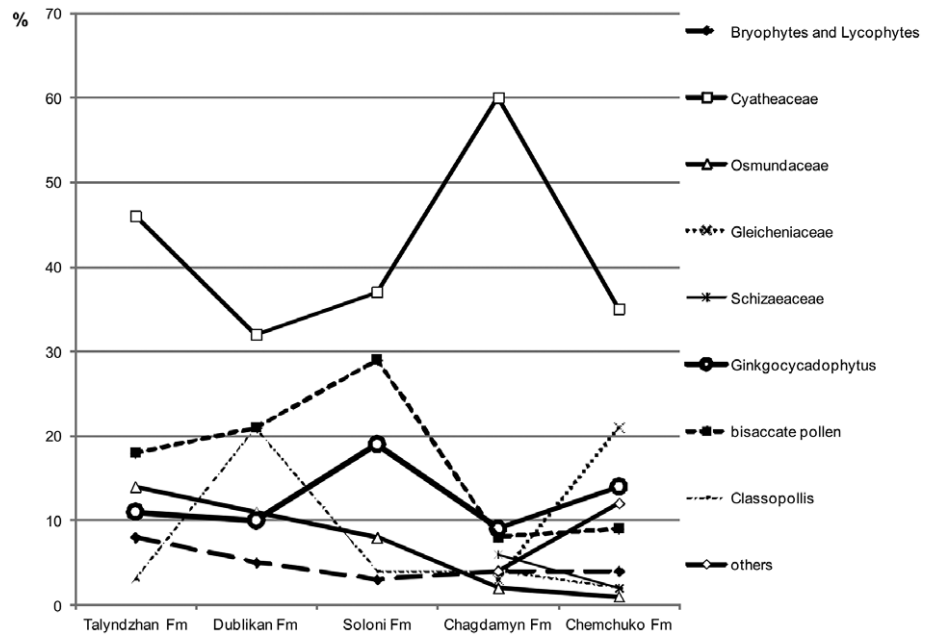


FIGURE 7 The relative abundances of major groups of spores and pollen in the coal-bearing deposits of the Bureya Basin

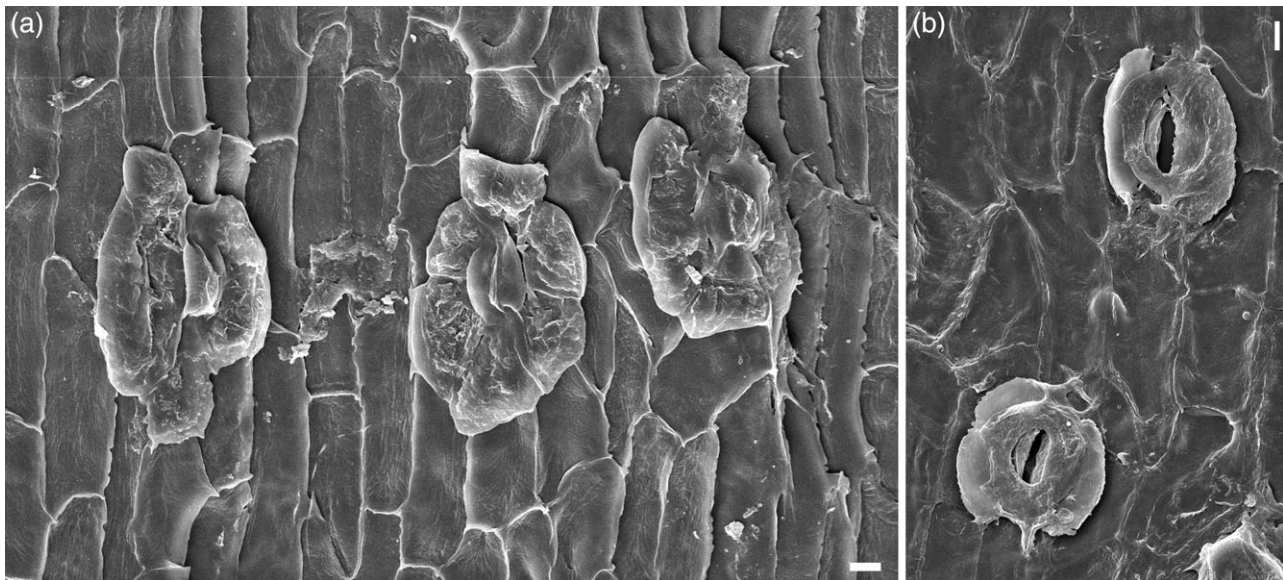


FIGURE 8 Cuticles of the coal-forming plants of the Bureya Basin. (a) *Bilsdalea* sp. (Dublikan Formation). (b) *Pseudotorellia angustifolia* Doludenko (Talyndzhan and Dublikan formations). Scale bars 10 μ m

increase of ferns and cheirolepidiaceous gymnosperms, which might be related to the marginal uplift and drier climate (Markevich, 1981, 1995).

The coal-forming plants are represented by *Pseudotorellia angustifolia* and conifer of unknown affinity *Bilsdalea* Harris, 1952 (Figure 8).

Palynological assemblage of the Soloni Formation (the Valangian) is dominated by conifers (up to 60 %) and *Ginkgocycadophytus* (Figure 6). Ferns reduce up to 50 %. The gleicheniaceous and schizaeaceous ferns have assumed great importance (see Figure 7).

In the Middle Cretaceous the sea retreated from this area and the coals are accumulated in the interior depressions. The palynoflora of the Chagdamyn Formation (the Barremian) is represented by mainly spores of *Cyathidites* (Figure 6). *Ginkgocycadophytus* and conifers prevail among gymnosperms. The diversity and amounts of ferns

are high in the palynological assemblage of the Chemchuko Formation (the Aptian); they are represented by cyatheaceous, gleicheniaceous, and osmundaceous ferns. The Pinaceae and Taxodiaceae predominate among gymnosperms. The involvement of *Ginkgocycadophytus* remains rather high (see Figure 7).

Krassilov (1972, 1973) pointed out the coal-forming role of *Sphenobaiera* and *Pseudotorellia longifolia* Doludenko in the Early Cretaceous of the Bureya Basin. The osmundaceous ferns lost their significance, gleicheniaceous and schizaeaceous ferns replaced them. Essential features of these plant communities are inherited from the Jurassic vegetation.

Based on a number of criteria, such as the abundance of bryophytes, cycadophyte index and replacement of ecological dominance, Krassilov suggested a warming trend during the deposition of sediments of Talyndzhan, Dublikan, and Chemchuko, a cooling trend in

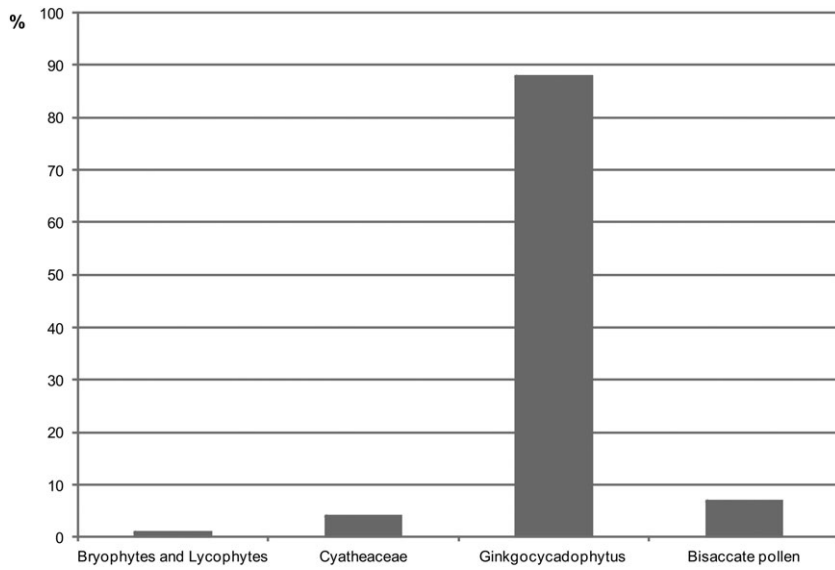


FIGURE 9 The relative abundances of major groups of spores and pollen in the coal seam of the Khilok Formation (Khilok Basin)

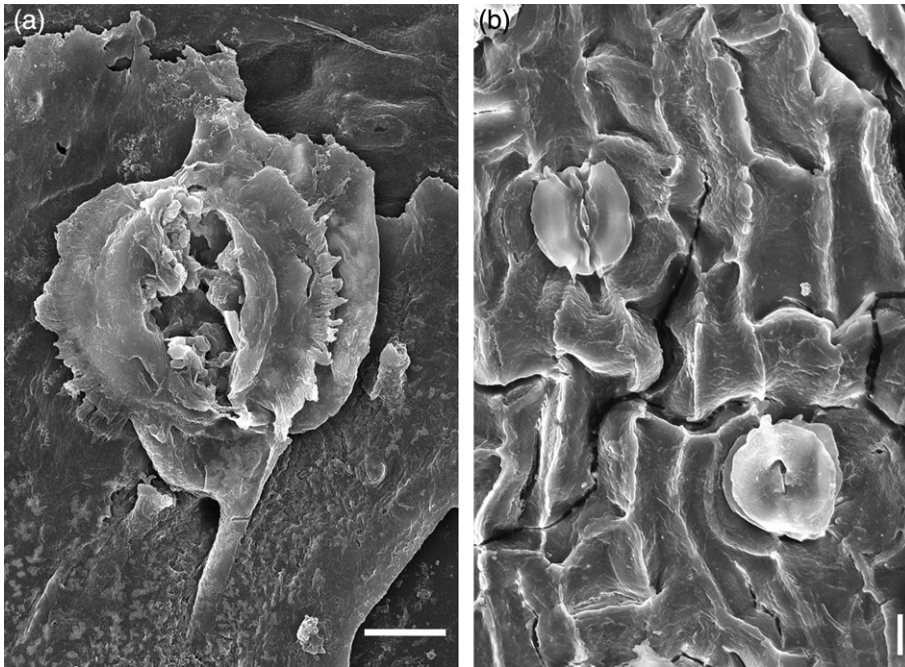


FIGURE 10 Cuticles of the coal-forming plants of the Doronino Formation (Chita-Ingoda Basin). (a) *Sphenobaiera starukhiniae* Bugdaeva. (b) *Pseudotorellia* sp. Scale bars 10 μ m

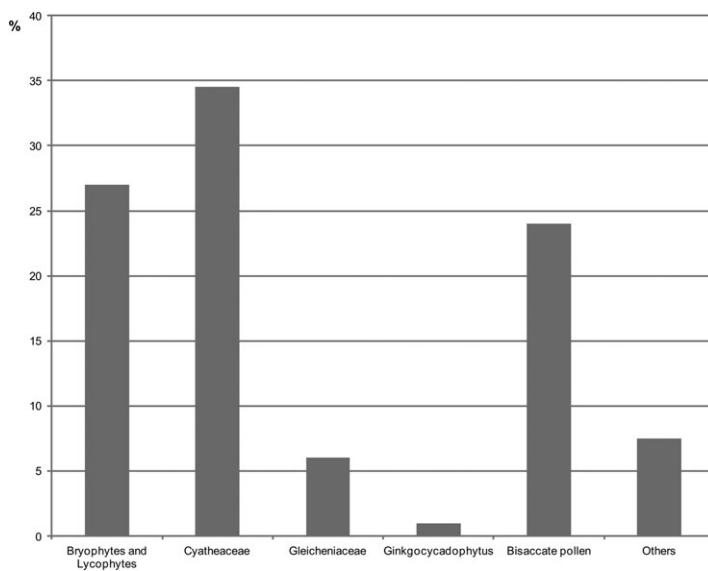


FIGURE 11 The relative abundances of major groups of spores and pollen in the coals of the Doronino Formation (Chita-Ingoda Basin)

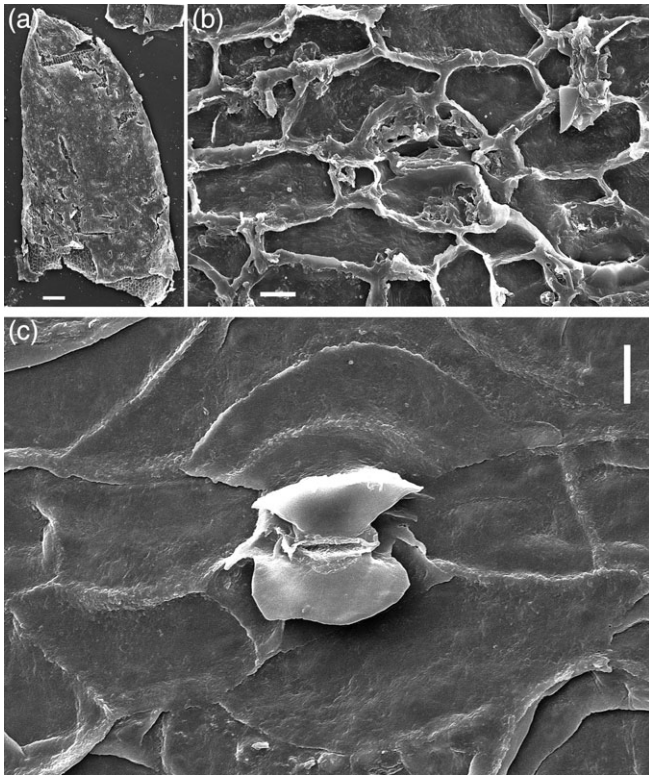


FIGURE 12 Cuticles of the coal-forming plants of the Doronino Formation (Bukachacha Basin) (a) Leaf of *Farndalea* sp., scale bar 200 μm . (b) Stoma of *Farndalea* sp., scale bar 10 μm . (c) *Pseudotorellia transbaikalica* Bugdaeva, scale bar 10 μm

Soloni and Chagdamin deposition. The czezanowskialeans are reduced, whereas cheirolepidiaceans increased during the Chemchuko climatic warming (Krassilov, 1973).

The cyatheaceous ferns, conifers *Bilsdalea*, ginkgoaleans (*Pseudotorellia angustifolia*, *P. longifolia*, *Sphenobaiera uralica* Krassilov, *S. ikorfatensis* (Seward) Florin) provided a basis for the Early Cretaceous mire plant communities.

3.2 | Transbaikalia

The Lower Cretaceous deposits fill numerous Transbaikalia intermontane rift basins. The strata are terrestrial, volcanogenic and coal-bearing (see Figure 2). The age of coal formation in this region was defined based on paleobotanical data (Bugdaeva, 1992, 1995; Bugdaeva & Markevich, 2012).

3.2.1 | Khilok Basin

The Khilok Formation (K_1hl) is represented by a sequence of trachyandesites, basalts, trachybasalts, their tuffs and tuffites with proluvial-lacustrine deposits (conglomerates, sandstones, and siltstones). Its geological age is the Barremian–Aptian (Gordienko, Bayanov, Klimuk, Ponomarchuk, & Travin, 1999).

In the sedimentary member of this sequence, E. V. Bugdaeva during the field expedition in 2002 found a thin coal seam (thickness is about 20 cm) consisting of very narrow dichotomously dissected leaves, later described as *Baierella averianovii* Bugdaeva. This representative of ginkgoaleans formed monospecific wetland plant communities and played coal-forming role during the Early Cretaceous (Bugdaeva & Markevich, 2007). The palynological spectrum from the coal seam is dominated by pollen *Ginkgocycadophytus* which was produced by ginkgoaleans (Figure 9).

3.2.2 | Chita-Ingoda Basin

In this basin the Barremian–Aptian Doronino Formation is widely distributed. This stratigraphic unit overlies with disconformity the Upper Jurassic strata and with conformity the Ulača or Byrtsa formations. The Doronino Formation consists of mainly siltstones, mudstones, and sandstones, trachyandesites, basalts, trachybasalts, their tuffs and tuffites prevail, and thick productive coal seams in the upper part (see Figure 2). The thickness of this formation is 200–1600 m.

The fossil flora of the Doronino Formation in this basin is dominated by ferns, ginkgoaleans, and conifers. It was revealed that the coals of this stratigraphic unit were mainly composed of the

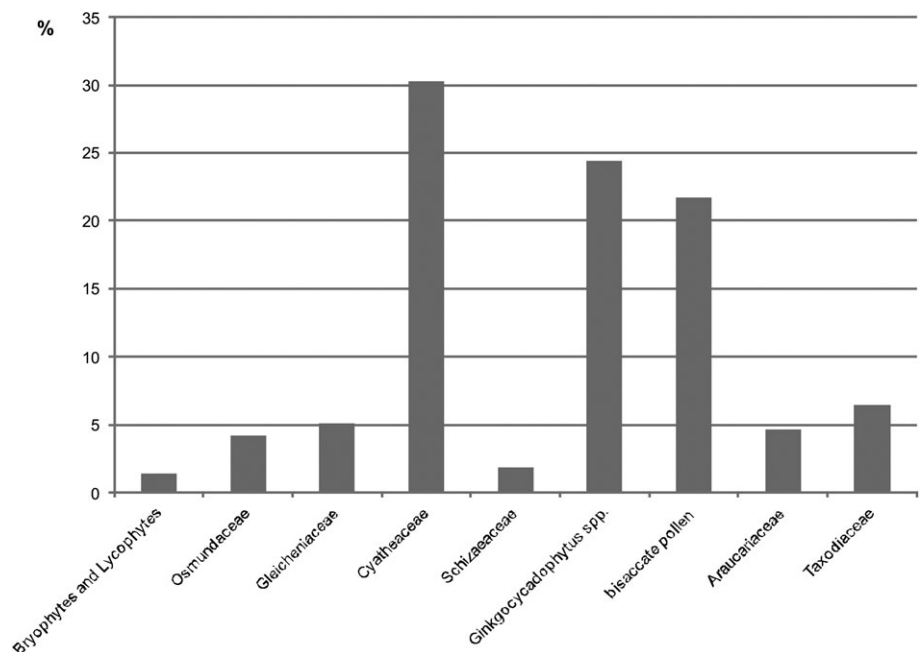


FIGURE 13 The relative abundances of major groups of spores and pollen in the coals of the Doronino Formation (Bukachacha Basin)

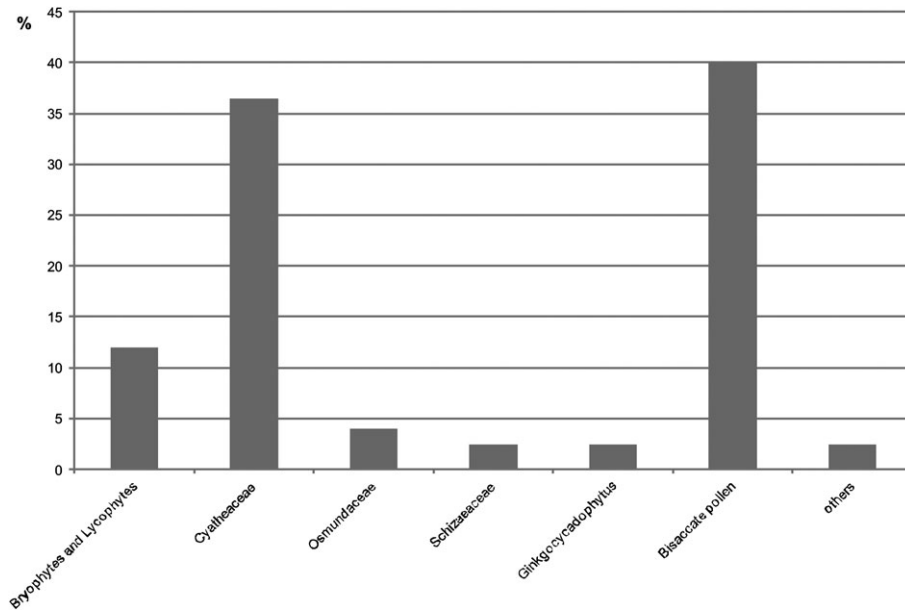


FIGURE 14 The relative abundances of major groups of spores and pollen in the coals of the Kuti Formation (Turga-Kharanor Basin)

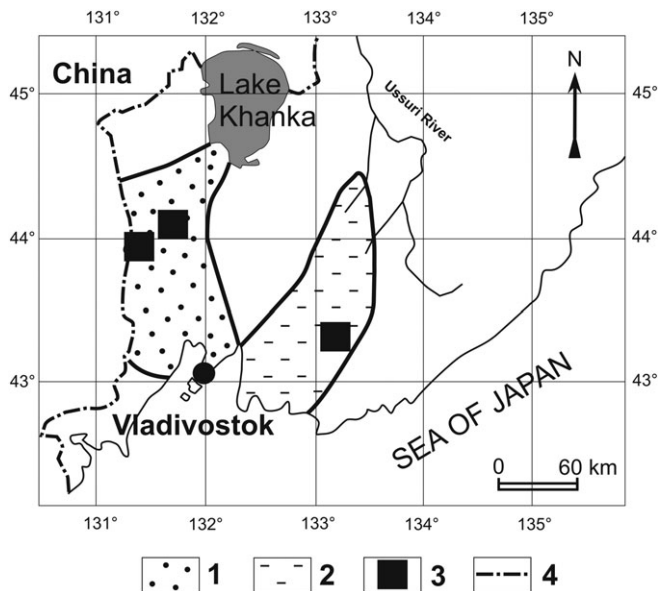


FIGURE 15 Study area in Primorye region. 1, Razdolnaya River Basin; 2, Partizansk Basin; 3, studied coal mines; 4, border between Russia and China

remains of ginkgoaleans, such as *Sphenobaiera starukhiniae* Bugdaeva (2010) and *Pseudotorellia* sp. (Figure 10a,b). The palynological spectra from coal seams are dominated by spores of bryophytes, lycophytes (Figure 6), and cyatheaceous ferns as well as by bisaccate pollen (Figure 11).

3.2.3 | Bukachacha Basin

The Doronino Formation in this basin is divided into two subformations. The lower subformation is represented by andesites, basalts, their lava breccias (150 m) and conglomerates, gravelstones, and coarse-grained sandstones (250 m). The upper subformation contains conglomerates, sandstones, mudstones, siltstones, and six productive coal seams, with a thickness of about 750 m (see Figure 2).

The fossil flora of coal-bearing deposits is dominated by conifers and ginkgoaleans, next in abundance—ferns and czekanowskialeans.

After bulk maceration of the coal from productive seam dispersed cuticles of *Farndalea* sp., belonging to conifers of araucariaceous affinity (Figure 12a,b), *Pityophyllum* sp. having pinaceous affinity, and ginkgoalean *Pseudotorellia transbaikalica* Bugdaeva (Figure 12c) were obtained.

The palynological spectrum of coal seam of the Doronino Formation in this basin is dominated by fern spores (more than 40%), *Ginkgo-cycadophytus* sp. (more than 26%) and bisaccate pollen having a pinaceous affinity (more than 22%). The percentage of araucariaceous pollen is about 4.5%, and that of taxodiaceous pollen 6.5%. Among the ferns, the representatives of Cyatheaceae dominate, the percentage of gleicheniaceous and osmundaceous spores (Figure 6) reaches 5% and 4%, respectively (Figure 13).

The paleobotanical and palynological data revealed that the main coal-forming plants during the Doronino deposition in the Bukachacha Basin were cyatheaceous ferns, extinct ginkgoalean *Pseudotorellia* and diverse conifers.

3.2.4 | Turga-Kharanor Basin

The sequence of this basin is terminated by the Kuti Formation, represented by alluvial and lacustrine sedimentary rocks. In the upper part this stratigraphic unit has thick productive coal seams (see Figure 2). The thickness is up to 1500 m.

The fossil flora of coal-bearing deposits is represented by ferns, conifers, and ginkgoaleans.

The cuticle of ginkgoalean *Pseudotorellia kharanorica* Bugdaeva and cheirolepidiaceous plant *Pagiophyllum* sp. was revealed from coal of the Kuti Formation (Bugdaeva, 1995).

The palynological spectra of coals are dominated by cyatheaceous fern spores and pinaceous pollen, in a less degree by bryophytes and lycophytes (Figure 14).

The main coal-forming plants in this basin were cyatheaceous ferns, conifers, and ginkgoaleans, probably bryophytes and lycophytes.

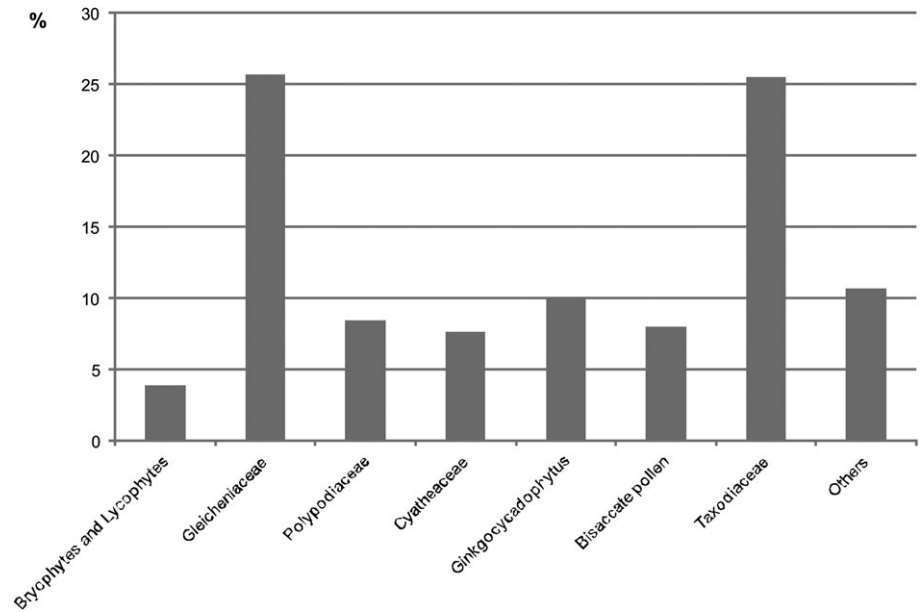


FIGURE 16 The relative abundances of major groups of spores and pollen in the coals of the Starosuchan Formation (Partizansk Basin)

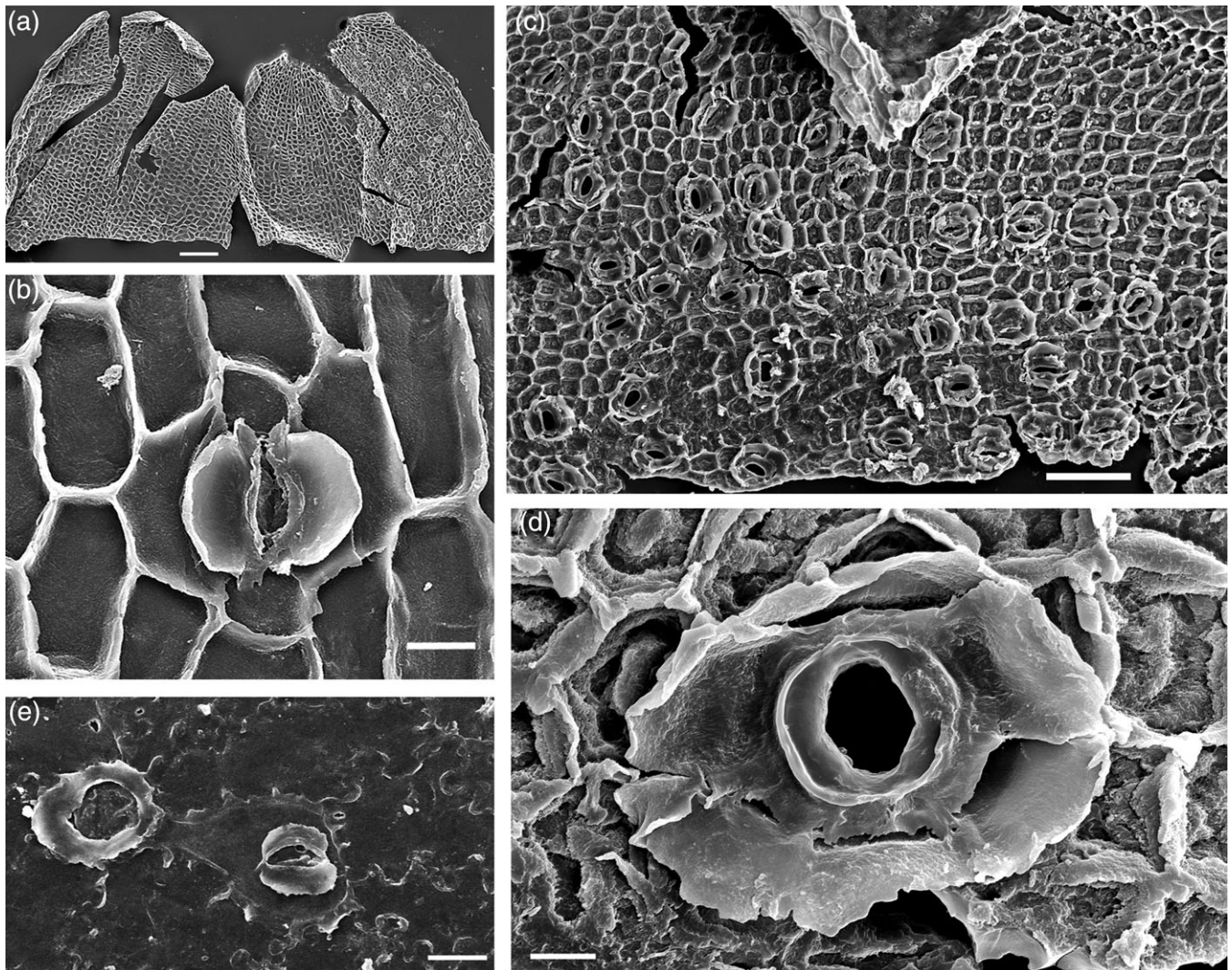


FIGURE 17 Cuticles of the coal-forming plants of the Lipovtsy Formation (Razdolnaya River Basin). (a) Upper part of leaf of *Mirovia orientalis* (Nosova) Nosova, scale bar 200 μm . (b) Stoma of *Mirovia orientalis*, scale bar 20 μm . (c) Stomatal and non-stomatal parts of *Athrotaxites orientalis* Deng et Chen, scale bar 100 μm . (d) Stoma of *Athrotaxites orientalis*, scale bar 10 μm . (e) Stoma and base of hair of *Nilssoniopteris rithidorachis* (Krysht.) Krassil., scale bar 20 μm

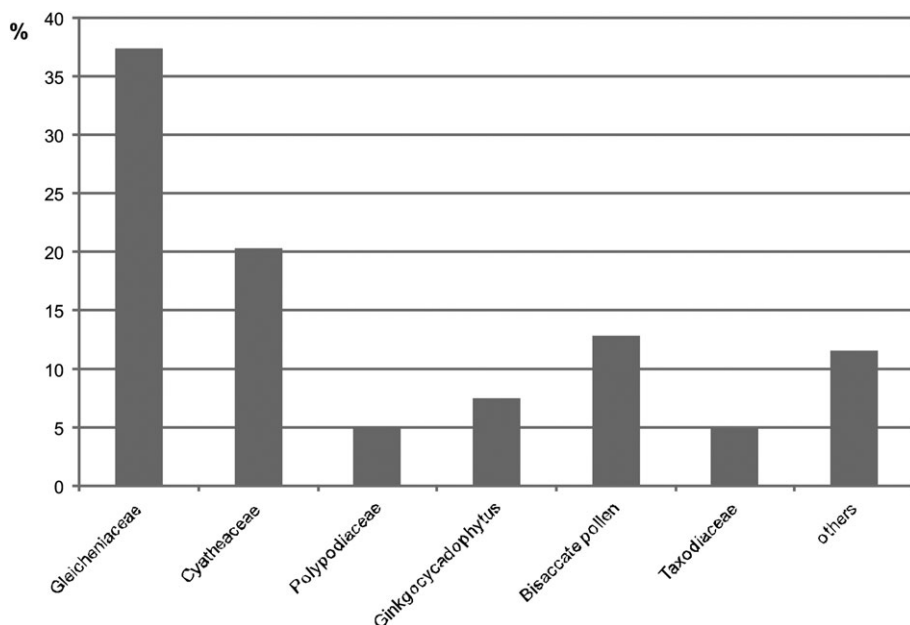


FIGURE 18 The relative abundances of major groups of spores and pollen in the coals of the Lipovtsy Formation (Razdolnaya River Basin)

3.3 | Primorye region

In the Primorye region two coal-bearing areas are widely distributed in the Partizansk and Razdolnaya basins (Figure 15).

3.3.1 | Partizansk Basin

The Starosuchan Formation of Aptian age was studied. The flora of this formation is dominated by the ferns and conifers. The coals of this stratigraphic unit are mostly composed of remains of taxodialean *Elatides asiatica* (Yok.) Krassil., subordinate Miroviaceae, rare ginkgoalean *Pseudotorellia* sp., and bennettite *Nilssoniopteris rithidorachis* (Krysht.) Krassil. (Bugdaeva, Markevich, & Volynets, 2014) The spores *Gleicheniidites* and pollen *Taxodiaceapollenites* (Figure 6) are dominants in the palynospectra (Figure 16).

3.3.2 | Razdolnaya River Basin

The unique resinous coals (rhabdopissites) occur in the upper part of the Aptian Lipovtsy Formation (see Figure 2). The flora of this stratigraphic unit is very abundant and diverse. The plants that produced coals mainly belong to the Miroviaceae (*Mirovia orientalis* (Nosova) Nosova), ginkgoalean *Pseudotorellia krassilovii* Bugd., to a lesser degree these coals are composed of other groups of conifers (including scale-leaved *Athrotaxites orientalis* Deng et Chen) and bennettites *Nilssoniopteris rithidorachis*, *N. prynadae* Samyl., and *Anomozamites* sp. (Figure 17). The palynological data have revealed that cyatheaceous and gleicheniaceus ferns have played an important role (Figure 18) in the formation of coal in this basin (Bugdaeva & Markevich, 2009).

4 | CONCLUSIONS

The Early Cretaceous coal-forming plants comprise cyatheaceous, gleicheniaceus, osmundaceous, and schizaeaceous ferns, ginkgoaleans (*Pseudotorellia*, *Sphenobaiera* and *Baierella*), bennettites, and

conifers having pinaceous, taxodiaceous, araucariaceous affinity, as well as extinct cheirolepidiaceus and miroviaceous plants.

It is well known that coal accumulation was controlled by tectonics, climatic conditions, and the structures of ecosystems that provided the organic material for the coal formation. The conservatism of taxonomic composition of mire plant communities is related to similar environments. The swamp plant communities and related coal accumulation seem to be sensitive to major environmental crises and these changes affect taxonomical composition of flora or extinction of swamp plants. A comparative account of the flora from different basins may substantiate the conclusions. Moreover palaeoecology of individual plant species may provide environmental inferences.

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