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Early Cretaceous changes of vegetation and environment in East Asia

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Abstract

Early Cretaceous floras of the Razdolnensky (Suifun) and Partizansky (Suchan) basins, currently located close to each other in South Primorye, differ significantly. The pre-Albian flora assemblages of the Razdolnensky basin are similar to those of northeastern China, whereas the pre-Albian assemblages of the Partizansky basin are close to the Rioseki complex of the Outer Japan. Tectonic reconstructions using these data have showed that the Voznesenka terrane, underlying Cretaceous sediments of the Razdolnensky basin, as well as the terranes of Inner Japan were displaced northwards along faults of the Tan-Lu system for more than 5° by latitude. The Sergeevka terrane, forming the basement of Early Cretaceous sediments of the Partizansky basin, as well as the terranes of Outer Japan, were displaced in the same direction for a much larger distance (about 15° by latitude). The major movements occurred in the earliest Albian as is suggested by the same Albian flora assemblages in the two basins. © 1999 Elsevier Science B.V. All rights reserved.

Keywords: Cretaceous; flora; palynoflora; East Asia; environment; tectonic reconstruction

1. Introduction

Cretaceous epicontinental sediments are abundant in the south of the Khanka superterrane (Primorye region) occurring in the Razdolnensky (Suifun) and Partizansky (Suchan) coal-bearing basins (Fig. 1). In the course of the study it has turned out that the Hauterivian–Albian flora assemblages of these basins are significantly different. Originally, this difference was explained by the presence at that time of a mountain structure separating the basins (Kras-

silov, 1967). A similar combination of Early Cretaceous floras was also found in Japan: heat-loving flora of the Ryoseki complex (Outer Japan) and cold-resisting flora of the Tetori complex (Inner Japan) (Kimura, 1987). Kimura considers this combination being due to tectonic displacements (Kimura, 1987).

We have analyzed specific compositions of the Early Cretaceous plants within the Partizansky and Razdolnensky basins, and compared these floras to those of northeast China and Japan. This paper represents, in brief terms, the results of our studies. On the basis of the data about the distribution of different types of the Early Cretaceous flora, and geological data, a new version of the pre-Albian reconstruction of the East Asia terrane positioning is offered.

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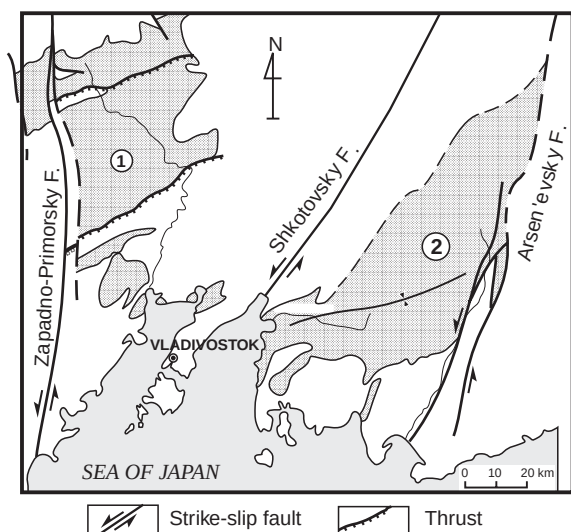


Fig. 1. Early Cretaceous epicontinental basins of South Sikhote-Alin: 1 = Razdolnensky (Suifun) basin; 2 = Partizansky (Suchan) basin.

2. Floral characteristics in basins of South Primorye

2.1. Partizansky (Suchan) basin

In the Partizansky (Suchan) basin, the Cretaceous section has at its base Berriasian volcanogenic and terrigenous sediments of the Monakinskaya Formation. The enclosed flora includes ferns, Cycadales, Bennettitales and Caytoniales. At the end of the Berriasian–Valanginian period (the Kapreevskaya and Klyuchevskaya Formations), the major wood-forming plants were various conifers associated with ferns and Cycadales. The fauna includes *Buchia*, *Ostrea* and *Corbicula*, as well as micro- and macroflora in the marine deposits of the Klyuchevskaya Formation (Tables 1 and 2).

During the Hauterivian and Barremian, forests were also dominated by conifers, but these were chiefly Taxodiaceae with an undergrowth consisting of ferns (including rare Schizaeaceae). The earliest of the Russian Far East flowering plants with the biporate pollen first appeared at the end of the Hau-

terivian (Table 1). It was likely connected with a transgression and related destabilization of the paleoenvironment.

Taxodiaceae remained dominant in the Aptian, but the role and diversity of the warmth- and water-requiring Gleicheniaceae increased among ferns. The angiosperms were represented by *Clavatipollennites*. This time was marked by active coal formation.

Terrigenous coal-bearing continental deposits formed in the basin from the Hauterivian to the middle Albian. During the Albian, the vegetation did not change markedly. However, among the conifers the typical Early Cretaceous species became extinct, while the diversity of angiosperms began to increase. A short-term transgression occurred in the middle Albian. At this time, the *Trigonia* layers were deposited. After that, the former continental sedimentation conditions reappeared.

The late Albian manifested a sharp change of vegetation in the basin. Diversity markedly decreased: numerous warmth-requiring species (*Gleicheniidites*, *Ginkgocycadophytus* and others) became extinct (Fig. 2). The niches released were occupied by pioneering plants, including angiosperms, mainly triporate pollen (Table 1).

Table 2 shows the distribution of plant megafossils of Tetori- and Ryoseki-types in the Early Cretaceous deposits of the Suchan and Suifun basins. It is apparent that these floras have mixed features.

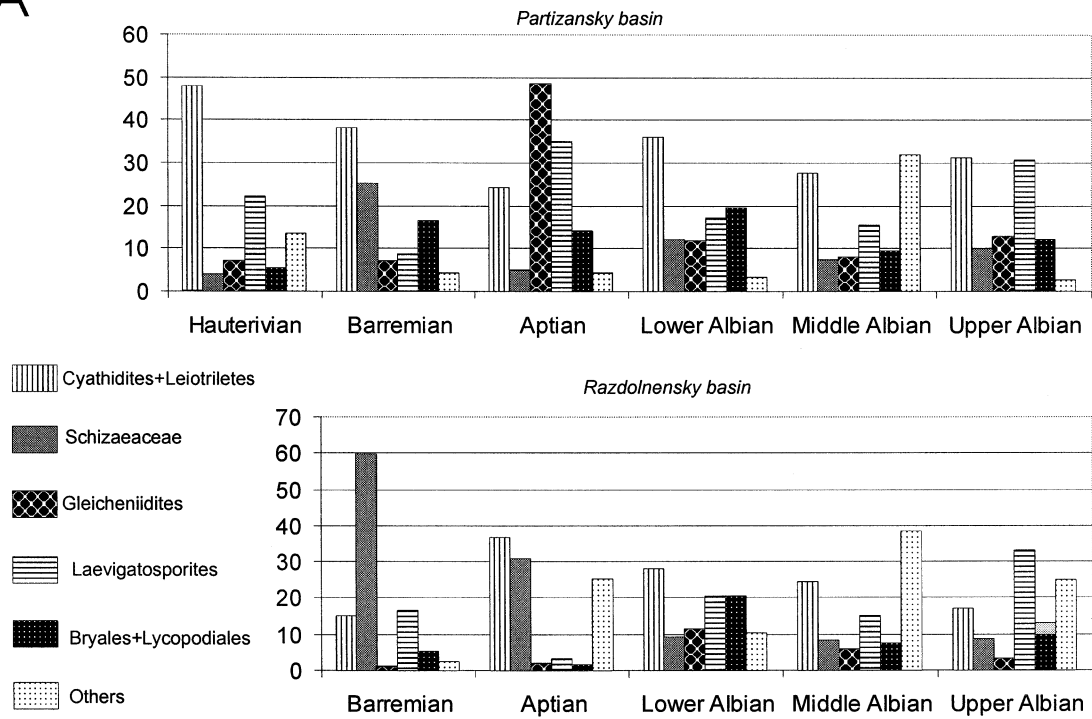
2.2. Razdolnensky (Suifun) basin

In the Razdolnensky (Suifun) basin, the base of the Cretaceous section contains Barremian deposits of the Ussuriyskaya Formation. In contrast to the Partizansky basin, the major wood-forming plants in the Barremian were various conifers (Podocarpaceae, Araucariaceae, Caytoniaceae with abundant *Ginkgocycadophytus*, Cyatheaceae and Schizaeaceae).

In the Aptian (the Lipovetskaya Formation), forests were dominated by the warmth- and water-requiring Gleicheniaceae, Matoniaceae and others. In the Albian, the diversity sharply decreased. Many of cycadophytes, conifers and ferns became extinct. The less

Fig. 2. The distribution of the fossil spores (A) and gymnosperm pollen (B) throughout the Lower Cretaceous in the South Sikhote-Alin region.

A



B

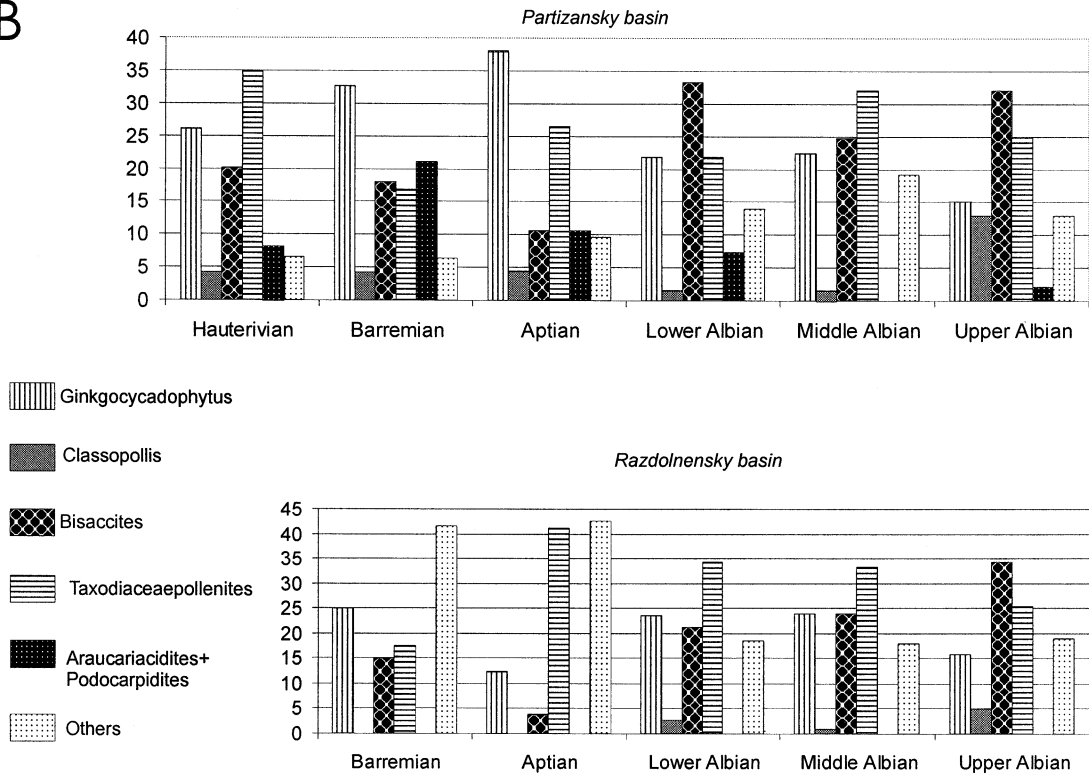


Table 1

Distribution of the fossil palynological taxa in the Early Cretaceous of northern China, Russia and Japan

Age	Taxa	North China	Russia		Japan	
			Suifun Basin	Suchan Basin	Hokkaido	Sea of Japan
Alb.	<i>Cyathidites (Leiotriletes)</i>	R	R	A		
	Schizaeaceae	C	C	C		R
	<i>Aequitriradites</i>	C	R			
	<i>Kuylisporites</i>	C	R			
	<i>Foraminisporites</i>	R	R			
	<i>Laevigatosporites</i>		A	A		C
	<i>Rouseisporites</i>	C				
	<i>Gleicheniidites</i>		R	C		R
	Taxodiaceapollenites			A		R
	Ginkgocycadophytus			C		
	<i>Classopollis</i>			C		A
	Bisaccate pollen			A		A
	Angiospermae	C	C	R		R
Apt.	<i>Cyathidites (Leiotriletes)</i>	R	A	A	A	C
	Schizaeaceae	A	C	R	C	C
	<i>Gleicheniidites</i>	R	A	A	C	
	<i>Osmundacidites</i>			C		
	<i>Rouseisporites</i>	R	R			
	<i>Kuylisporites</i>	R	R			R
	<i>Foraminisporites</i>					R
	<i>Classopollis</i>	R	R	C	A	A
	Inaperturopollenites	C	C	A	C	
	Bisaccate pollen	A	A	C		
	Angiospermae	C			R	
Brm.	<i>Cyathidites (Leiotriletes)</i>	C	C	A		
	Schizaeaceae	A	A	S		
Hau.	<i>Gleicheniidites</i>	C	C	C		
	<i>Aequitriradites</i>	R	R	R		
	<i>Neoraistrickia</i>	A	S	S		
	<i>Classopollis</i>	R	R	R		
	Bisaccate pollen	A	A	C		
	Inaperturopollenites	R	R	A		
	Angiospermae			R		
Vlg.	<i>Cyathidites (Leiotriletes)</i>	A		A		
	Schizaeaceae	R		R		
Ber.	Ginkgocycadophytus	R		C		
	<i>Araucariacidites</i>	R		C		
	<i>Classopollis</i>	R		A		
	Bisaccate pollen	A		A		

A = abundant; C = common; R = rare.

warmth-requiring *Laevigatosporites*, Bryales, *Lycopodiumsporites*, *Equisetites* became dominant.

The difference in vegetation between the Partizansky and Razdolnensky basins disappeared in the Albian (Fig. 2). Since this time, plants of these two basins had many common characteristics.

We will discuss causes of the different fossil palynoflora compositions of the basins close to each other in space and time (Table 1) and compare them with the simultaneous floras of Japan (Takahashi, 1974; Markevich, 1981, 1995) and northeastern China (Song, 1988; Wang et al., 1990; Li, 1994).

Table 2

Distribution of fossil plants of Tetori- and Ryoseki-types in the Early Cretaceous of South Sikhote-Alin region (Suchan and Suifun basins)

Age	Fossil plant taxa	Suifun basin	Suchan basin
Alb.	<i>Coniopteris burejensis</i> (Zal.) Sew.	R	R
	<i>Neozamites denticulatus</i> (Krysht. et Pryn.)	T	
	<i>Nilssonia</i> ex gr. <i>brongniartii</i> (Mant.) Dunk.	R	
	<i>Nilssonia densinervis</i> -type	R	
	<i>Ginkgo pluripartita</i> (Schimp.) Heer	T	T–R
	<i>Sphenobaiera</i> sp.	T	
	<i>Podozamites tenuinervis</i> Heer	T	
Apt.	<i>Elatides asiatica</i> (Yok.) Krassil.	T	T
	<i>Nathorstia pectinata</i> (Goepp.) Krassil.	R	
	<i>Matonidium goeppertii</i> (Ett.) Schenk.	R	
	<i>Weichselia reticulata</i> (Stokes et Webb) Ward	R	
	<i>Alsophilites nipponensis</i> (Oishi) Krassil.	R	
	<i>Coniopteris burejensis</i> (Zal.) Sew.	R	R
	<i>Zamites borealis</i> Heer	R	
	<i>Dictyozamites grossinervis</i> Yok.	T	
	<i>D. cordatus</i> (Krysht.) Pryn.	T	
	<i>Ctenis latiloba</i> Krysht. et Pryn.	T	
	<i>C. yokoyamae</i> Krysht.	T	
	<i>C. intermedia</i> (Krysht. et Pryn.) Pryn.	T	
	<i>Nilssonia densinervis</i> -type	R	R
	<i>Nilssonia nicanica</i> Pryn.	R	
	<i>N.</i> ex gr. <i>brongniartii</i> (Mant.) Dunk.	R	R
	<i>Ginkgo pluripartita</i> (Schimp.) Heer	T	
	<i>Podozamites</i> ex gr. <i>lanceolatus</i> (Lindl. et Hut.)	T	T
	<i>P. tenuinervis</i> Heer	T	
	Araucariaceae	R	
	<i>Elatides asiatica</i> (Yok.) Krassil.	T	T
Bar.	<i>Osmunda</i> sp.	T	
	<i>Coniopteris burejensis</i> (Zal.) Sew.	R	R
	<i>Alsophilites nipponensis</i> (Oishi) Krassil.		R
	<i>Neozamites denticulatus</i> (Krysht. et Pryn.)	T	T
	<i>Dictyozamites cordatus</i> (Krysht.) Pryn.	T	
	<i>Ptilophyllum bajulae</i> Krassil.	R	
	<i>Nilssonia densinervis</i> (Font.) Berry	R	
	<i>N.</i> ex gr. <i>brongniartii</i> (Mant.) Dunk.	R	R
	<i>Ginkgo pluripartita</i> (Schimp.) Heer		T
	<i>Pseudotorellia</i> sp.	T	
	<i>Podozamites</i> ex gr. <i>lanceolatus</i> (Lind. et Hutt.)	T	T
	Araucariaceae	R	R
Hau.	<i>Alsophilites nipponensis</i> (Oishi) Krassil.		R
Vlg.	<i>Coniopteris burejensis</i> (Zal.) Sew.		R
	<i>Coniopteris arctica</i> (Pryn.) Samyl.		R
	<i>Dictyozamites kawasaki</i> Tateiwa		T
	<i>Podozamites</i> ex gr. <i>lanceolatus</i> (Lind. et Hutt.)		T
Ber.	<i>Alsophilites nipponensis</i> (Oishi) Krassil.	R	R
	<i>Coniopteris arctica</i> (Pryn.) Samyl.		R
	<i>C. burejensis</i> (Zal.) Sew.	R	R
	<i>C.</i> cf. <i>vsevolodii</i> E. Lebed.	R	
	<i>C. setacea</i> Pryn.	R	
	<i>Dictyozamites falcatus</i> (Morris)	T	T

Table 2 (continued)

Age	Fossil plant taxa	Suifun basin	Suchan basin
Ber.	<i>Nilssonia densinervis</i> -type		R
	<i>N. pseudomediana</i> Dobrusk.		R
	<i>N. ex gr. brongniarti</i> (Mant.) Dunk.		R
	<i>Pseudotorellia</i> sp.		T
	<i>Brachyphyllum</i> aff. <i>expansum</i> (Sternb.) Sew.		T

R = representative of Ryoseki-type plants; T = representative of Tetori-type plants.

The Early Cretaceous floras of northern China and the Suifun basin are related. Floras of Outer Japan and the Suchan basin are closely analogous.

3. Tectonic reconstruction

When making a reconstruction, account must be taken of the large-scaled lateral displacements that took place in Cretaceous and Miocene times in this region (Xu et al., 1989; Xu, 1993; Otofujii and Matsuda, 1993; Otofujii et al., 1996).

The map of pre-Miocene tectonic structures of East Asia (Fig. 3) was made using the models of Kropotkin and Shakhvarstova (1965) and Bersenev et al. (1987) and the paleomagnetic data of Otofujii and Matsuda (1993) and Otofujii et al. (1996). The system of continental-margin shear zones is shown after Xu et al. (1989), Utkin (1993) and Xu (1993). The distribution of different-type Early Cretaceous floras and palynological assemblages is given using the data of Krassilov (1967), Kimura (1987), Ohana and Kimura (1995) and Markevich (1995).

Tectonic reconstruction of the eastern margin of Asia for the Late Albian time (Fig. 3) consisted in recovering the original position of the terranes which had been displaced northward along the continental-margin left-lateral faults during the Early Cretaceous (Golozoubov and Khanchuk, 1996). The reconstruction suggests that the Jurassic and Lower Cretaceous terranes were always situated in the vicinity of the continent–ocean border, while they were being formed and transported. This is evidenced by the arcotic composition of most of the Triassic, Jurassic and Lower Cretaceous terrigenous formations in Sikhote-Alin (Markevich, 1985). It is also suggested that the saw-shapes northeastern margins of the terranes resulted from left-lateral replacements along

the shear zones. The similarities in structure and history of certain terranes in Sikhote-Alin and Japan — specifically, the Samarka terrane of Sikhote-Alin and the terranes Mino, Tamba, Ashio, Northern Kitakami and others of Outer Japan, the Taukha terrane of Sikhote-Alin and the Southern Chichibu terrane of Outer Japan, the Sergeevka terrane of Sikhote-Alin and the terranes Abakuma and Southern Kitakami of Japan — were taken into account as well. In addition, the reconstruction was based on the distribution of the different-type Early Cretaceous floras on the continental margin.

4. Conclusions

A main conclusion is that the Jurassic and Lower Cretaceous terranes of Outer Japan and southeastern Primorye have been markedly (not less than 15° of latitude) displaced northward from the place of their formation (15–30°N). These terranes chiefly are fragments of the Middle Jurassic–Early Cretaceous subduction-related melange that may be traced as an almost continuous belt along the southeastern margin of Asia.

The terranes of southwestern Primorye and Inner Japan have drifted in the same direction, but not so far (5° of latitude).

Distribution of the different-type Early Cretaceous mega- and palynofloras on the continental margins of Eastern Asia allows a more precise evaluation of the Early Cretaceous left-lateral displacements. The paleomagnetic data (Hirooka, 1990) and data on the distribution of the different-type Early Cretaceous malacofauna (Kononov, 1988; Matsukawa et al., 1993) are less certain, but do not contradict the reconstruction presented.

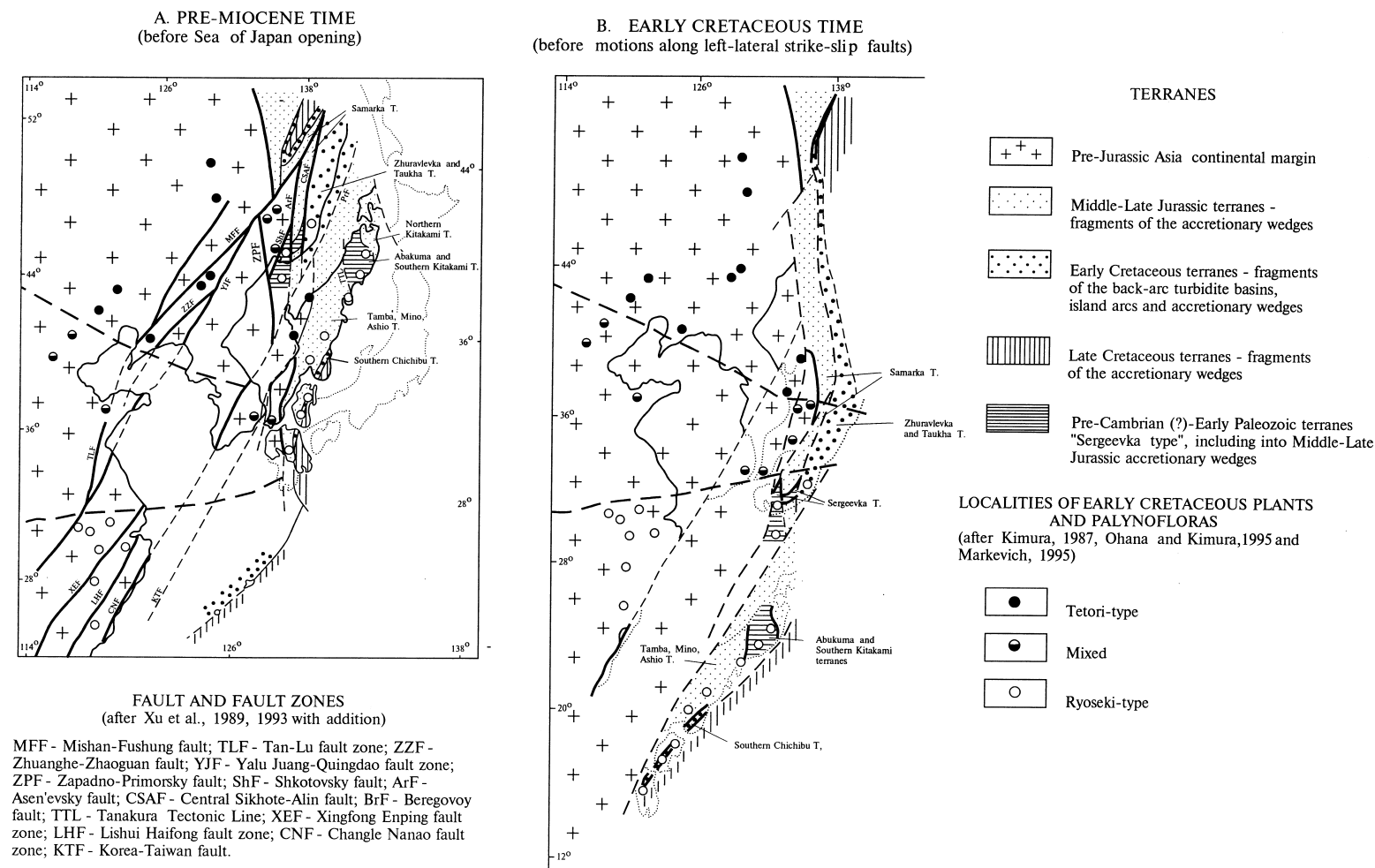


Fig. 3. Tectonic reconstruction of the eastern margin of Asia for the Pre-Miocene (A) and the Early Cretaceous (B).

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